



## Baltic Sea acidification

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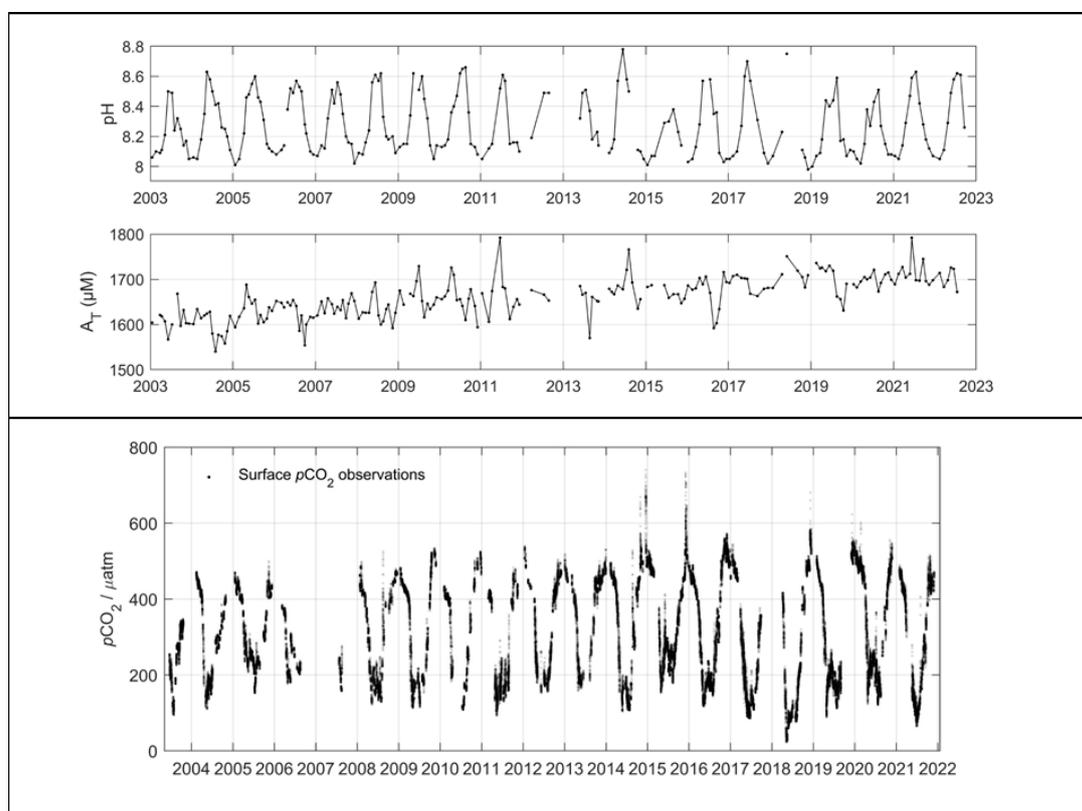
## 1 Key message

The Baltic Sea carbonate system is undergoing large changes due to: 1) increasing CO<sub>2</sub> in the atmosphere, 2) increasing inputs of total alkalinity (A<sub>T</sub>) from land, 3) changes in the balance between primary production and respiration due to eutrophication/oligotrophication, and 4) warming, shifting the carbonate speciation towards higher CO<sub>2</sub> partial pressure (pCO<sub>2</sub>) and lower pH levels.

In some areas of the Baltic Sea, pH is decreasing and pCO<sub>2</sub> is increasing at rates substantially higher than experienced from ocean acidification alone.

Although the impact of current levels of pH and pCO<sub>2</sub> on populations of indigenous marine organisms is small, continued acidification may threaten species that cannot adapt to a future warmer and acidified Baltic Sea. Organisms exposed to upwelling of hypoxic, corrosive water from the deeper basins are particularly affected.

A<sub>T</sub> and pH are monitored by the SMHI on a monthly basis, whereas pCO<sub>2</sub> is monitored continuously on the commercial ship Finnmaid in a cooperation between IOW and SYKE (Figure 1). Such observations can be used to estimate long-term trends as well as seasonal variations of these key variables. While longer time series exist for the open Baltic Sea, many other areas only have a few spot samples, particularly in the coastal zone. This means that a broader understanding of the acidification development is limited by a sparse spatial and temporal monitoring of key variables.



**Figure 1.** Upper: pH and total alkalinity (A<sub>T</sub>) data from the monitoring station BY15 (Eastern Gotland Basin) provided by the SMHI. Lower: pCO<sub>2</sub> data from an area close to BY15 on the shipping route of Finnmaid.

## 1.1 Citation

The data and resulting data products (e.g. tables, figures and maps) available on the indicator web page can be used freely given that it is used appropriately and the source is cited. The indicator should be cited as follows:

HELCOM (2023) Baltic Sea acidification. HELCOM element indicator report. Online. [Date Viewed], [Web link].

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## 2 Relevance of the indicator

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Ocean acidification reduces calcification of shell forming organisms and affects a range of physiological processes especially related to cellular ion regulation. Based on current knowledge, most organisms react negatively to acidification but the responses vary between species and communities. Primary producers may benefit from the improved carbon uptake energetics, whereas negative effects are most common in macrozoobenthos, but potentially also in fish. Acidification may lead to shifts in plankton community composition and alter flows of energy and organic matter partitioning. The Baltic Sea plankton system may shift towards the microbial loop (lower trophic level heterotrophy), with negative effects on food supply to the higher trophic levels. The effects of acidification are strongly modified by interactions with other drivers, including increases in temperature, stratification, deoxygenation and eutro-/oligotrophication.

### 2.1 Ecological relevance

Ocean acidification threatens calcifying organisms, including clams, mussels and brittle stars (Dupont and Pörtner 2013, Gao et al. 2019). It leads to decreased saturation states of the mineral forms of calcium carbonate, aragonite and calcite, which calcifying organisms generally use in their skeletal structures. Acidification also alters various other physiological processes, especially ones related to cellular ion regulation, and the non-calcifying organisms, including invertebrates and fish, are often equally adversely affected (Thor and Dupont, 2018). Photoautotrophic organisms may also benefit from rising pCO<sub>2</sub> levels through improved carbon uptake energetics. Growth of macroalgae generally increases with rising pCO<sub>2</sub> (Gao et al., 2019), whereas phytoplankton responses range between communities and species from positive to negative (Gao et al., 2019; Hutchins and Fu, 2017).

Most information on the species responses comes from relatively short-term incubations, and there are major gaps in the knowledge on effects from long-term exposure to acidification (Pansch et al. 2018), and the capacity of species and ecosystems to adapt to rising pCO<sub>2</sub> (Riebesell and Gattuso, 2015; Vargas et al., 2017). While attention largely has focused on single-species response to a single factor (Riebesell and Gattuso, 2015), recent studies emphasize the need to account for the effects on communities and habitat structure and complexity (Riebesell and Gattuso, 2015; Sunday et al., 2017), and on multiple stressors including acidification, global warming, increased stratification and deoxygenation (Jutterström et al., 2014; Riebesell and Gattuso, 2015; Boyd et al., 2018). Warming, for example, intensifies stratification during the warm season leading to increased nutrient limitation, which in some areas might override the effect of increased CO<sub>2</sub> supply on phytoplankton (Hutchins and Fu, 2017).

In the Baltic Sea, biota naturally faces large fluctuations in pH, mainly driven by diurnal and seasonal decoupling of primary production and respiration, with ranges far exceeding the atmospheric signal predicted for the next 100 years (e.g. Rossoll et al., 2013). Especially in coastal areas, frequent upwelling events entrain CO<sub>2</sub> rich deep water into the surface layer, drawing the surface water pH further down. It is likely that long term acclimation

and adaptation of species (Frommel et al., 2013; Pansch et al. 2013a) and adaptation of the communities (Rossoll et al., 2013) to high pCO<sub>2</sub> and strong pH fluctuations occur in the Baltic Sea. However, peak pCO<sub>2</sub> values in productive estuarine areas could increase to values >4000 µatm, most probably having detrimental effects on e.g. calcifying organisms (Thomsen et al., 2010).

#### Plankton communities

The Baltic Sea plankton communities appear tolerant to pCO<sub>2</sub> levels up to 1000-1400 µatm (Rossoll et al., 2013; Wulff et al., 2018). Phytoplankton and nitrogen fixing cyanobacteria show mainly subtle or no response to acidification (e.g. Kremp et al., 2012; Karlberg and Wulff, 2013; Paul et al., 2015; 2016; 2018). Higher pCO<sub>2</sub> levels may affect the plankton community composition e.g. by supporting picoeukaryotic primary producers and small sized microzooplankton and decreasing the diversity of microzooplankton community in summer (Crawford et al., 2017; Paul et al., 2015). The mesozooplankton community was found tolerant to acidification in a mesocosm experiment (Lischka et al., 2017). The same was true for offspring production of the copepod *Eurytemora affinis* (Almén et al., 2016), whereas development of copepod *Acartia tonsa* nauplii was slower at high pCO<sub>2</sub> levels (Vehmaa et al., 2016). There is no available information on the response of gelatinous zooplankton in the Baltic Sea, but the results from nearby areas indicate they respond neutrally or positively to acidification (Havenhand et al., 2019).

Acidification may alter the energy transfer and organic matter partitioning in the planktonic systems. Net primary production has been demonstrated to increase with increasing pCO<sub>2</sub> in mesocosm experiment, but without concomitant increase in C export, implying that increased primary production would not counteract the increasing atmospheric CO<sub>2</sub> concentration (Paul et al., 2015; Spilling et al., 2016). Some studies indicate that the toxin production of toxic dinoflagellate blooms may increase with acidification (Kremp et al., 2012).

#### Benthic communities

Calcifying organisms, such as bivalve *Limecola balthica* (previously *Macoma balthica*) and blue mussels of genus *Mytilus*, belong to key species of the benthic communities in the Baltic Sea. All life forms of *L. balthica* have been shown to react negatively to acidification (Jansson, 2017). Even a small decrease in pH may slow down the growth rate and survival of the larval stage (Jansson, 2017), and metabolic rates and energy demand of adult stage increase with acidification (Jakubowska and Normant-Saremba, 2015; Jansson, 2017). Adult bivalves *Mytilus edulis* and *Arctica islandica* appear robust against acidification until 1400-1700 µatm (Hiebenthal et al., 2013; Thomsen et al., 2010; 2013). Calcification and shell growth of *M. edulis* however may decrease with increasing acidification (Thomsen and Melzner, 2010), and extreme pCO<sub>2</sub> levels of > 4000 µatm most likely affect the shell growth (Thomsen et al., 2010; Thomsen and Melzner, 2010). Non-calcifying larval stages of barnacles are generally robust to near future acidification levels (Pansch et al., 2013b). The growth and survival of adult barnacles differs notably between traits, with higher tolerance to high pCO<sub>2</sub> levels in barnacles from habitats with high pCO<sub>2</sub> fluctuations

compared to barnacles from more stable pCO<sub>2</sub> habitats (Pansch et al., 2014). Sensitive populations have been shown to only partially acclimate to long-term elevated pCO<sub>2</sub>, with significant effects on reproduction (Pansch et al., 2018).

Negative impacts of acidification on adult bivalve growth largely occur through increasing energy demand for maintaining the physiological balance (Jansson, 2017; Thomsen and Melzner, 2010). It is suggested that sufficient food supply would mediate the ability of calcifying species to cope with rising pCO<sub>2</sub> levels (Thomsen et al., 2010; Pansch et al., 2014). Effects of acidification on mussel calcification can also be counteracted by macrophyte-driven elevation of mean pH and temporal refuge from acidification due to fluctuations in the carbonate system in dense algae and seagrass habitats (Wahl et al., 2018). Growth of seagrass *Zostera marina* slightly benefits from the rising pCO<sub>2</sub> concentration (Eklöf et al., 2012; Takolander et al., 2017), whereas responses of bladder wrack *Fucus vesiculosus* differ between studies from negative to weakly positive (Takolander et al., 2017; Graiff et al., 2015). A review of the current knowledge indicates that red algae respond positively to acidification, whereas for green algae both positive and negative responses have been recorded (Takolander et al., 2017). The community level response of microvegetation to acidification further depends on the relative growth responses of *Fucus vesiculosus* and *Zostera marina* and their epiphytes (Wahl et al., 2018).

## Fish

Commercially and ecologically important fish species, cod and herring, potentially respond negatively to acidification in the Baltic Sea. The fertilization success of cod in laboratory experiments was unaffected by pCO<sub>2</sub> <1360 µatm (Frommel et al., 2010), and egg stages of both herring and cod have been found robust to high levels of pCO<sub>2</sub> (Franke and Clemmesen, 2011; Frommel et al., 2013). Response of larval cod to acidification experiments has ranged from tolerance to high pCO<sub>2</sub> levels (~3200 µatm; Frommel et al., 2013) to doubling of daily mortality at pCO<sub>2</sub> ~1100 µatm (Stiasny et al., 2016). Acidification may reduce the protein synthesis of the herring embryos implying reduction of larval growth (Franke and Clemmesen, 2011). Fish populations may be further hampered by the effects of acidification on the food web: a review of the current knowledge indicates that acidification drives the energy flow towards microbial loop, reducing the energy supply of higher trophic levels, zooplankton and fish (Havenhand et al., 2019).

## 2.2 Policy relevance

Increasing CO<sub>2</sub> in the atmosphere is the major driver for acidification of marine waters. The 2021 HELCOM Baltic Sea Action Plan reiterated that the effects of climate change on the Baltic Sea are already evident, and that climate change will continue to have an increasingly significant impact on the Baltic Sea ecosystem, necessitating even more stringent action. The BSAP emphasized the need for continued research and adaptive management to mitigate the effects and strengthen the resilience of the Baltic Sea to climate change by reducing other human pressures on the ecosystems and also underscored the need to further adapt HELCOM's policies and Recommendations to take into account the effects of climate change. The BSAP also includes action HT5 on

developing a strategic approach to ocean acidification for the Baltic Sea with first steps addressing the knowledge gaps by 2025.

The indicator also addresses the following qualitative descriptors and criteria of the MSFD for determining good environmental status:

1. D1C6 The condition of the habitat type, including its biotic and abiotic structure and its functions (e.g. its typical species composition and their relative abundance, absence of particularly sensitive or fragile species or species providing key function, size structure of species), is not adversely affected due to anthropogenic pressures.
2. D6C5 The extent of adverse effects from anthropogenic pressures on the condition of the habitat type, including alteration to its abiotic and biotic functions (e.g. its typical species composition and their relative abundance, absence of particularly sensitive or fragile species or species providing key function, size structure of species), does not exceed a specified proportion of the natural extent of the habitat type in the assessment area.

The indicator further addresses MSFD D5 “Human-induced eutrophication is minimised, especially adverse effects thereof, such as losses in biodiversity, ecosystem degradation, harmful algae blooms and oxygen deficiency in bottom waters”. The relationship of the indicator with eutrophication is dual: Firstly, eutrophication intensifies pCO<sub>2</sub> fluctuations and leads to seasonal and subsurface acidification, and secondly, acidification may benefit primary production and hence intensify eutrophication.

The indicator is also relevant for United Nations’ Sustainable Development Goal 14 Target “Minimize and address the impacts of ocean acidification, including through enhanced scientific cooperation at all levels”. The policy relevance is indicated in Table 1.

**Table 1.** Policy relevance of the indicator.

	<b>Baltic Sea Action Plan (BSAP)</b>	<b>Marine Strategy Framework Directive (MSFD)</b>
<b>Fundamental link</b>	Segment: Horizontal topics, theme: Climate change  <ul style="list-style-type: none"> <li>• Actions and overview relevant to multiple topics and processes.</li> </ul>	Descriptor 1 Pelagic habitat.  <ul style="list-style-type: none"> <li>• Criteria 6 The condition of the habitat type, including its biotic and abiotic structure and its functions (e.g. its typical species composition and their relative abundance, absence of particularly sensitive or fragile species or species providing a key function, size structure of species), is not adversely affected due to anthropogenic pressures.</li> <li>• Feature – Pelagic broad habitats.</li> </ul>

		<ul style="list-style-type: none"> <li>• Element of the feature assessed – Pelagic broad habitat types.</li> </ul>
<b>Complementary link</b>	<p>Segment: Biodiversity</p> <p>Goal: “Baltic Sea ecosystem is healthy and resilient”</p> <ul style="list-style-type: none"> <li>• Ecological objective: “Functional, healthy and resilient food webs”.</li> <li>• Management objective: ”Reduce or prevent human pressures that lead to imbalance in the food web”.</li> </ul> <p>Segment: Eutrophication</p> <p>Goal: “Baltic Sea unaffected by eutrophication”</p> <ul style="list-style-type: none"> <li>• Ecological objective: “Natural distribution and occurrence of plants and animals”.</li> </ul>	<p>Descriptor 5 Eutrophication.</p> <ul style="list-style-type: none"> <li>• Human-induced eutrophication is minimised, especially adverse effects thereof, such as losses in biodiversity, ecosystem degradation, harmful algae blooms and oxygen deficiency in bottom waters.</li> </ul> <p>Descriptor 7 Hydrographical conditions.</p> <ul style="list-style-type: none"> <li>• Criteria 1 Spatial extent and distribution of permanent alteration of hydrographical conditions (e.g. changes in wave action, currents, salinity, temperature) to the seabed and water column, associated in particular with physical loss of the natural seabed.</li> <li>• Feature – Broad habitats.</li> </ul>
<b>Other relevant legislation:</b>	<p>1. EU Nitrates Directive; EU Urban Waste-Water Treatment Directive; Industrial Emissions Directive (IED), Water Framework Directive, WFD; the Gothenburg Protocol to Abate Acidification, Eutrophication and Ground-level Ozone under UNECE Convention on Long-range Transboundary Air pollution (CLRTAP); ); EU NEC Directive (2016/2284/EU); Water Code of Russian Federation; Federal Act on the internal maritime waters, territorial sea and contiguous zone of the Russian Federation; IMO designated the Baltic Sea as a ”special area” for passenger ships under MARPOL (International Convention for the Prevention of Pollution from Ships) Annex IV (on sewage from ships); EC Directive 2000/59/EC on port reception facilities; NOx emission control area (NECA) in the Baltic and North seas designated by IMO.</p>	

### 2.3 Relevance for other assessments

At present, acidification is not considered as part of the holistic assessment but the topic and report will be used to develop key interactions with relevant topics such as eutrophication or pelagic habitats under the relevant thematic assessments.

### 3 Threshold values

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There are no universal thresholds for acidification parameters that apply to all organisms, but two main aspects of acidification should be considered, hypercapnia and decalcification. It should also be mentioned that increasing  $p\text{CO}_2$  may stimulate productivity, thereby favouring some species over others, which may alter the community compositions and food-web interactions (Havenhand et al., 2019). However, our current knowledge on such indirect effects is limited and will not be considered here.

Hypercapnia is the elevation of  $p\text{CO}_2$  in the environment, which may cause intracellular pH to decrease, potentially reaching critical levels where essential physiological cell functions cease, or may increase energy costs of maintaining relatively constant intracellular pH, changing the fitness and competition between organisms and thereby altering communities. As indicated in Section 2.1, these responses are variable among species and even among stages, e.g. juvenile stages appear to be more sensitive to hypercapnia. Another important issue is the length of exposure to high  $p\text{CO}_2$  levels. Whereas most organisms do not exhibit direct pathological effects to elevated  $p\text{CO}_2$  levels in their present acidification regime, the overall shifting baseline of decreasing pH and increasing  $p\text{CO}_2$  may potentially lead to detrimental acidification levels precipitating more acute physiological effects. However, establishing critical thresholds of  $p\text{CO}_2$  for Baltic Sea marine species poses a scientific challenge that needs to be resolved.

Acidification alters the speciation of the carbonate system, leading to lower the saturation states for the two calcium carbonate polymorphs that marine organisms use to construct their shells or tests – aragonite and calcite (Doney et al., 2009). The saturation state for these two forms are expressed by their  $\Omega$ -values, with values  $<1$  indicating undersaturation, which may cause decalcification of calcifying organisms. Except for the Gulf of Bothnia, basins in the Baltic Sea are typically oversaturated in calcium carbonate during the productive period, when production upregulate pH, albeit most basins also display undersaturated conditions during winter. Most marine organisms are still capable of calcifying even if  $\Omega < 1$  due to the organic matrix of their skeletons and tests (Melzner et al., 2013). Consequently,  $\Omega = 1$  cannot be considered a critical threshold for decalcification. Nevertheless, calcifying organisms in waters with undersaturated conditions do not exploit their full growth potential, e.g. shells from blue mussels are smaller in waters with low pH and  $\Omega$  –values less than 1 (Melzner et al., 2011; Gazeau et al., 2010). As for  $p\text{CO}_2$ , establishing critical thresholds for  $\Omega$ -values is a scientific challenge that should be addressed.

#### 3.1 Setting the threshold value(s)

As indicated in Section 3, setting threshold values for  $p\text{CO}_2$  and  $\Omega$  remains a scientific challenge and is not part of the current indicator approach. An elements indicator is an indicator that chronicles an important processes or factor of direct relevance to the marine environment. There are no threshold values applied in this indicator at this stage and potentially threshold values may not be directly applicable to an Element indicator.

The Element indicator acts dominantly as a fact sheet to which other relevant topics can reference.

## 4 Results and discussion

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### 4.1 Status assessment

Marine ecosystems exhibit variations in parameters of the carbonate system on daily, seasonal and interannual scales (Duarte et al., 2013), with ranges exceeding 1 pH unit in poorly buffered systems like parts of the Baltic Sea (Carstensen and Duarte, 2019). Such variations can be natural and organisms have adapted to them. Similarly, there are large gradients in the carbonate system parameters across the Baltic Sea, where pH in the surface gradually increases from mean levels around 7.7 in the poorly buffered Bothnian Bay to 8.1 in the Baltic Proper towards the Kattegat (Table 2). On average, the Bothnian Bay and the Bothnian Sea are presumably sources of atmospheric CO<sub>2</sub>, whereas the remaining parts presumably operate as CO<sub>2</sub> sinks. pH is typically higher during summer when production dominates over respiration, leading to a net uptake of CO<sub>2</sub>, and lower during winter when respiration dominates over production, leading to a net release of CO<sub>2</sub>. The seasonal variability in pH is largest in the Gulf of Finland (range = 0.84) and lowest in the Kattegat (range = 0.17), whereas seasonal variability in pCO<sub>2</sub> is largest in the Baltic Proper (range = 363 μatm) and lowest in the Kattegat (range = 148 μatm). In bottom waters where respiration dominates, pH values are lower than in surface waters and pH below 7 is not uncommon in oxygen depleted zones. Correspondingly, pCO<sub>2</sub> typically increases with depth, reaching as high as 3-4000 μatm in the Landsort Deep and Gotland Basin.

**Table 2.** Mean level and seasonal variability, given as range of monthly means, in surface pH and pCO<sub>2</sub> (in  $\mu\text{atm}$ ) assessed by the BALTSEM model (Gustafsson and Gustafsson 2020) for a period with low anthropogenic influence, compared to more recent conditions (2000-2017) assessed from monitoring data of pH and A<sub>T</sub> at representative HELCOM stations. Note that not all basins had sufficient data to determine all the parameters of the carbonate system and that model results aggregated to less resolution than the HELCOM Assessment units (as indicated by footnotes) reflecting the sparse spatial coverage of the observations.

Basin	Historical conditions (1850-1900)				Contemporary conditions			
	Mean		Seasonal range		Mean		Seasonal range	
	pH	pCO <sub>2</sub>	pH	pCO <sub>2</sub>	pH	pCO <sub>2</sub>	pH	pCO <sub>2</sub>
Bothnian Bay <sup>1</sup>	7.88	351	7.73-7.97	279-503	7.71	454	7.62-7.92	284-545
Bothnian Sea <sup>2</sup>	8.06	307	7.96-8.18	225-360	7.96	383	7.69-8.24	173-630
Baltic Proper <sup>3</sup>	8.15	303	8.09-8.27	218-369	8.13	350	7.91-8.42	165-528
Gulf of Finland <sup>of</sup>	8.07	349	7.87-8.30	193-537	8.07		7.67-8.51	
Gulf of Riga	8.20	345	8.07-8.31	250-483				
Bornholm Basin	8.16	296	8.09-8.27	215-376	8.10	375	7.93-8.28	236-519
Arkona Basin	8.16	293	8.08-8.27	215-370	8.08	345	7.92-8.44	249-445
Danish Straits <sup>4</sup>	8.14	291	8.04-8.31	181-378				
Kattegat	8.13	292	8.06-8.32	179-359	8.07	359	7.98-8.15	284-432

<sup>1</sup>Includes Bothnian Bay and Northern half of The Quark

<sup>2</sup>Includes Bothnian Sea, Southern half of The Quark and Åland Sea

<sup>3</sup>Includes Eastern and Western Gotland Basin, and Northern Baltic Proper

<sup>4</sup>Includes The Sound, Great Belt, Kiel Bay and Bay of Mecklenburg

Thus, there are natural spatial gradients in pH and other carbonate system parameters across the Baltic Sea. These gradients are mainly linked to salinity gradients, but also to differences in productivity and thus CO<sub>2</sub> uptake by autotrophs. Temporal variations are on the other hand predominantly controlled by the metabolic balance between production

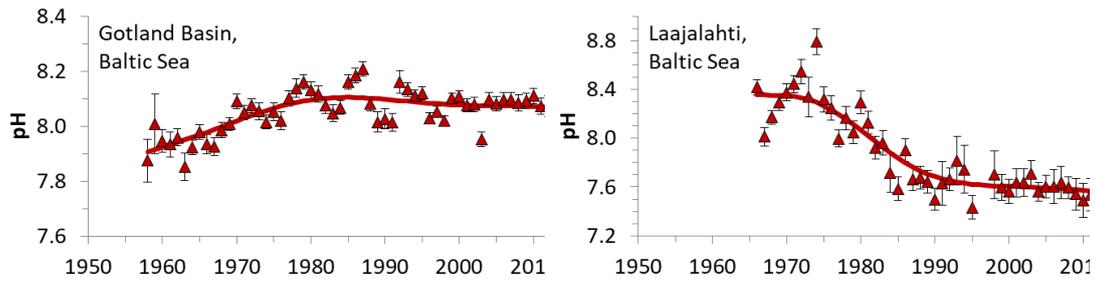
and respiration, although hydrology also plays a role in oligotrophic systems (Carstensen and Duarte, 2019). Temporal variability in pH increases with eutrophication in high-latitude systems, where production and respiration are seasonally decoupled, and in stratified systems, where production and respiration further are spatially decoupled. Thus, in order to assess the ecosystem effect of acidification, it is necessary to consider deviations from the natural mean level and variability, and to combine these deviations with assessments of potential consequences for the naturally occurring biota as well as their interactions with repercussions for ecosystem functioning (Havenhand et al., 2019).

#### *Added value as eutrophication indicator*

Apart from the fact that carbonate system parameters, in particular pH, provide the most direct indicator for acidification, monitoring of the carbonate system also provides an additional powerful indicator of eutrophication. A generally accepted definition of the term “eutrophication” was given by Nixon (1995) who considers eutrophication as “an increase in the rate of supply of organic matter to an ecosystem”. Regarding the autochthonous supply, hence, biomass production, observations of the surface water carbonate system constitute an ideal tool to quantify organic matter production rates and thus monitor eutrophication. This was demonstrated by Schneider and Müller (2018) who, for the base case year 2009, used high resolution surface water pCO<sub>2</sub> records to calculate the net community production. The latter corresponds to the amount of organic matter that has the potential to cause oxygen depletion and hydrogen sulphide formation after sinking into deeper water layers. The production rates determined for the Baltic Sea proper on the basis of the pCO<sub>2</sub> data exceeded the estimates obtained from the traditional nutrient approach by up to 70 %. The reasons for this discrepancy are not fully understood, but can partly be explained by non-Redfield production stoichiometry, in particular during the mid-summer cyanobacteria bloom. It has to be emphasized that the oxygen demand at depth resulting from organic matter decomposition is mainly linked to the oxidation of carbon, and thus carbon-based productivity estimates provide the closest link to oxygen deficiency.

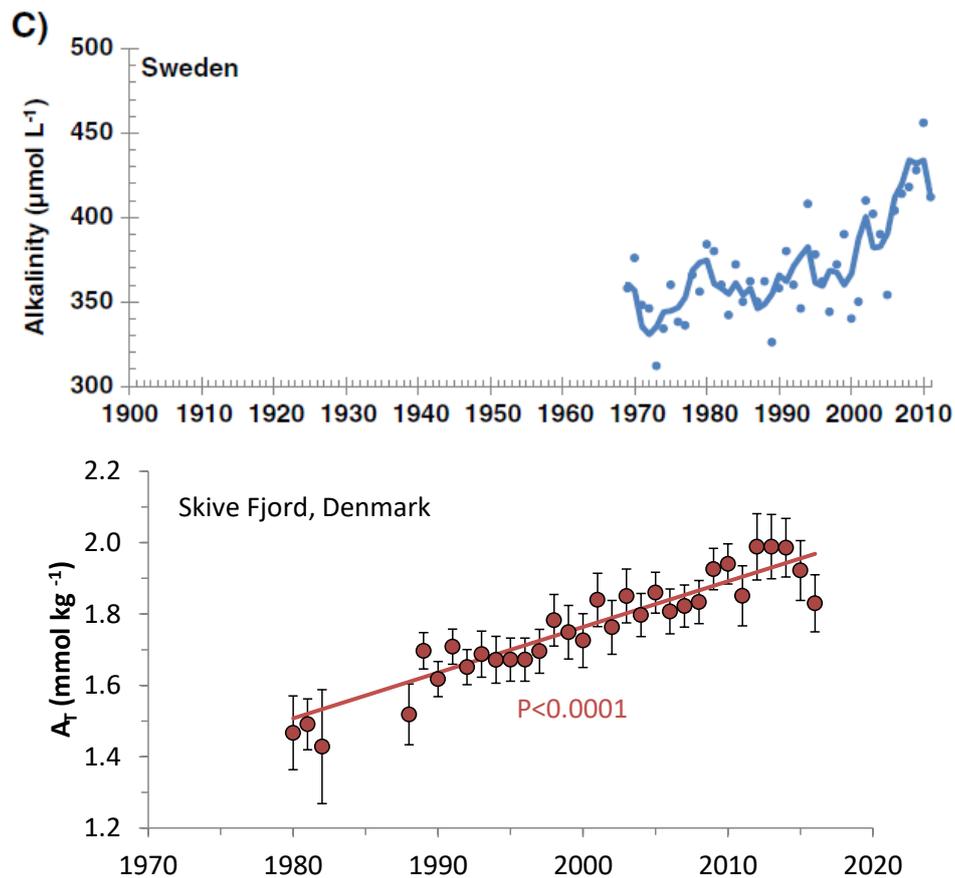
## 4.2 Trends

Ocean acidification due to increasing atmospheric CO<sub>2</sub> concentrations causes a predictable decline of pH (~0.02 per decade) and increase of pCO<sub>2</sub> (~2 µatm per year), whereas coastal systems experience changes over time which can exceed those of the ocean significantly (Duarte et al., 2013; Carstensen and Duarte, 2019). The Baltic Sea is no exception to this (Figure 2), where first eutrophication apparently led to increasing pH in surface waters due to enhanced uptake of CO<sub>2</sub> and later nutrient reduction caused pH to decline. Thus, pH changes in the Baltic Sea are tightly coupled to eutrophication/oligotrophication.



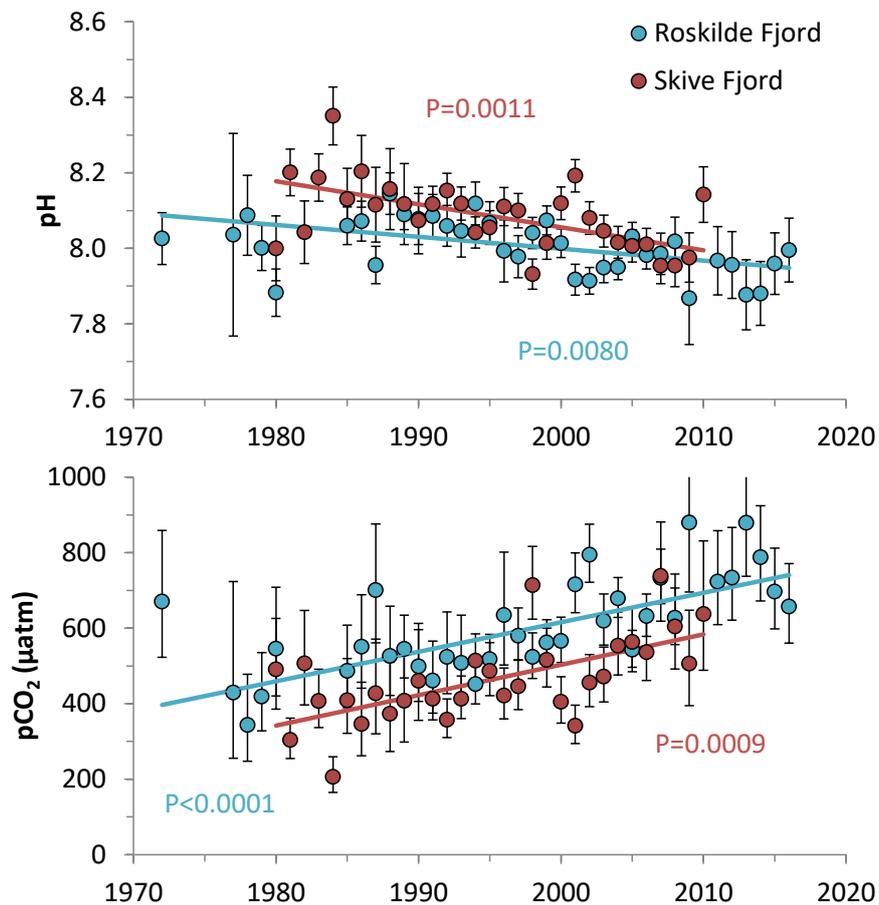
**Figure 2:** Trends in pH estimated from monitoring data from the Gotland Basin and Laajalahti Bay near Helsinki. From Carstensen and Duarte (2019).

Another process influencing the carbonate system is the change in  $A_T$ .  $A_T$  inputs change with land use and enhanced weathering from climate change (Raymond and Cole, 2003). This is also observed for the Baltic Sea (Figure 3).



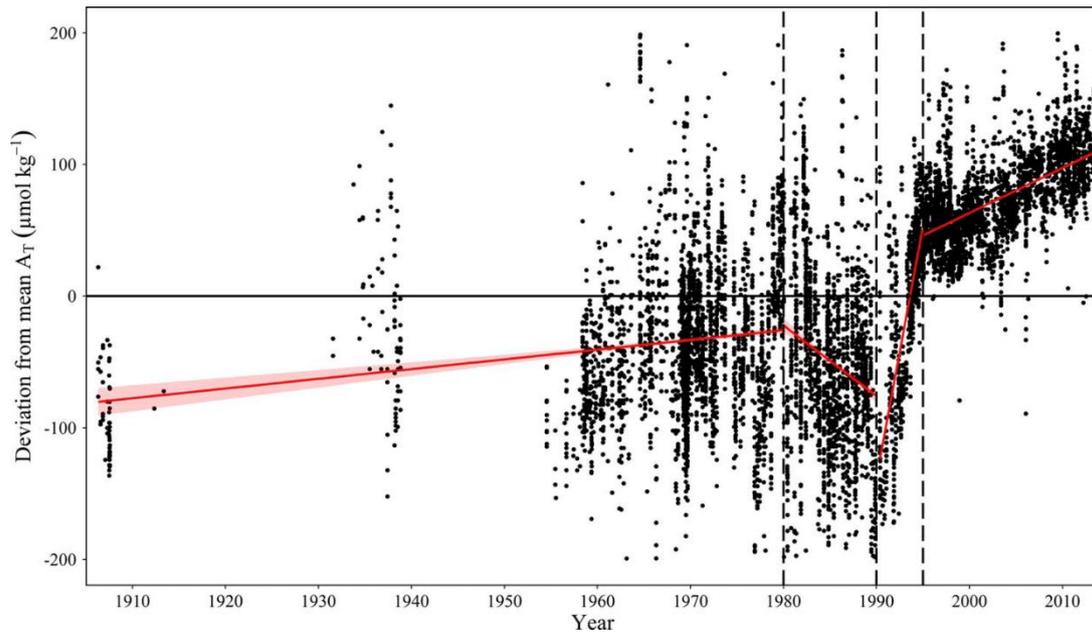
**Figure 3:** Flow-weighted concentration of  $A_T$  in Swedish rivers ([www.slu.se](http://www.slu.se)) and Danish streams discharging into Skive Fjord. From Duarte et al. (2013) and Carstensen et al. (2018).

Despite the increased  $A_T$  buffering acidification, pH is decreasing and  $pCO_2$  is increasing significantly over time in some Danish estuaries (Figure 4). The pH declines are ~2.5 times faster than ocean acidification (OA) and caused by the combination of increasing atmospheric  $CO_2$  (~1 OA trend), reduced nutrient input (~1 OA trend) and warming (~0.5 OA trend) (Carstensen et al., 2018). However, it should be stressed that these trends do not uniformly apply to all coastal areas and the open Baltic Sea.



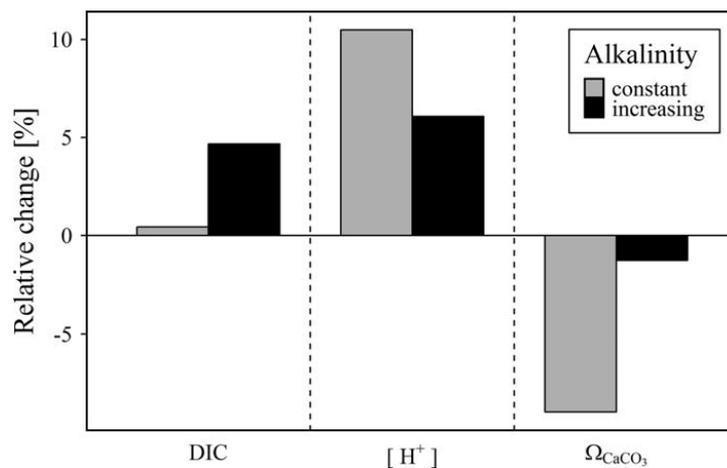
**Figure 4:** Trends in pH and  $pCO_2$  in two Danish estuaries. From Carstensen et al. (2018).

For the major basins, trends in  $A_T$  have recently been statistically evaluated based on a compilation of all available  $A_T$ -data, including the CANIBAL data, the SMHI data base, the BONUS Baltic-C data compilation and the FMI monitoring data (Müller et al., 2016). The evaluation shows that  $A_T$  measurements since 1995, where standard reference materials for  $A_T$  and dissolved inorganic carbon ( $C_T$ ) became available (Dickson et al., 2007), can be very reliably used for a trend analysis. The authors show an increase of  $A_T$  of  $\sim 70 \mu mol kg^{-1}$  between 1995 and 2014, corresponding to an annual increase of  $3.4 \mu mol kg^{-1} yr^{-1}$  (Figure 5). The highest trends of  $7.4 \mu mol kg^{-1} yr^{-1}$  was observed in the Gulf of Bothnia at a salinity of 3, which represents an increase of almost 20% over the same time frame.



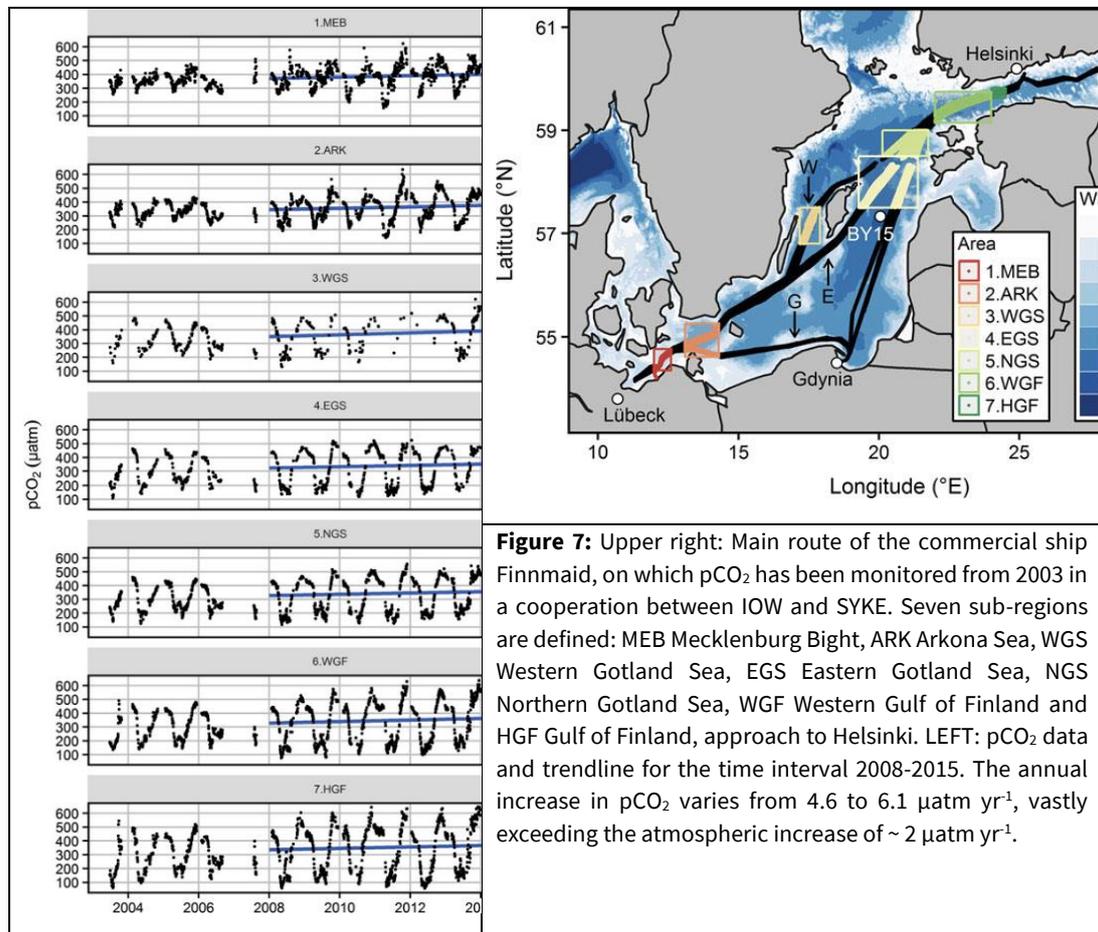
**Figure 5:** Temporal  $A_T$  trends in the Central Baltic Sea from 1900 to 2015. Displayed are deviations ( $d_{AT}$ ) of observed  $A_T$  from mean  $A_T$  values for individual salinity intervals of 0.2 (over a range from 6.5-7.7). From Müller et al. (2016).

Müller et al. (2016) further demonstrate that for the central Baltic Sea, the observed  $A_T$  trend has the potential to largely compensate for the pH decrease due to rising atmospheric  $CO_2$  levels during the same time frame, to increase the storage potential for carbon dioxide, and to almost entirely mitigate the reduction of carbonate saturation ( $\Omega_{CaCO_3}$ ), (Figure 6).



**Figure 6:** Combined influence of simultaneous changes in atmospheric  $pCO_2$  and seawater  $A_T$  on dissolved inorganic carbon (DIC), proton concentration  $[H^+]$ , and calcium carbonate saturation ( $\Omega_{CaCO_3}$ ) in the Central Baltic Sea from 1995 to 2014. Based on an atmospheric  $pCO_2$  increase from 360 to 400  $\mu atm$ , the relative changes in DIC,  $[H^+]$ , and  $\Omega_{CaCO_3}$  were computed for two  $A_T$  scenarios: The observed  $A_T$  trend of  $3.4 \mu mol kg^{-1} yr^{-1}$  was taken into account (black bars) and the  $A_T$  was assumed to be constant at mean  $A_T$  (grey bars). From Müller et al. (2016).

Due to the larger trend of up to  $7.4 \mu\text{mol kg}^{-1}$ , these effects are even stronger in the Northern Basins, in particular the Bothnian Bay where the increase in  $A_T$  could completely compensate for acidification. However, with the lack of other carbonate system parameters measured simultaneously, the real effect of the observed  $A_T$  trend cannot be addressed.



For the assessment of the carbonate system in surface waters, the measurement of  $p\text{CO}_2$  through water-air equilibration coupled to infrared spectroscopy developed into an internationally used approach (Dickson et al., 2007), which led to the SOCAT database for surface water  $p\text{CO}_2$  (Bakker et al., 2016). Continuous measurements have been recorded on the commercial vessel Finnmaid since 2003 and recently been compiled in Schneider and Müller (2018). Surface water  $p\text{CO}_2$  trends amounted  $4.6$ - $6.1 \mu\text{atm yr}^{-1}$  in the individual basins, exceeding the atmosphere  $p\text{CO}_2$  increase (approximately  $2 \mu\text{atm yr}^{-1}$ ) by a factor 2-3 (Figure 7), with reasons currently being unclear. This shows the complexity of the processes that control the carbonate system and highlights the necessity to monitor pH. Surface monitoring of  $p\text{CO}_2$  is currently initiated on several other commercial vessels as action of the BONUS Project BONUS INTEGRAL, see <https://www.io-warnemuende.de/integral-home.html>.

### 4.3 Discussion text

The Baltic Sea is experiencing acidification to a variable degree depending on location. Observed carbonate system parameters display trends over decadal time-scales – trends that depend on several different factors such as the increasing atmospheric CO<sub>2</sub> level, changes coupled to eutrophication/oligotrophication, as well as changes in A<sub>T</sub>. The impact of acidification in the Baltic Sea is so far small, but unless CO<sub>2</sub> emissions are reduced, the potential future effects on the marine ecosystem are considerable (Section 2.1). Our knowledge on this environmental problem is however scattered among few sites, where data are available and studies have been made.

## 5 Confidence

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The confidence in assessing status and trends of carbonate system parameters depends on the spatial and temporal coverage of data as well as the accuracy of measurements. Measurement of carbonate system parameters is straightforward, although precise pH measurements is a challenge (see Section 9.2). Long-term time-series of pH and  $A_T$  are available for several open water stations and a few coastal sites, but the limited coverage of monitoring data both spatially and temporarily affects the confidence in assessing status and trends of acidification development in the Baltic Sea on a broader scale.

The confidence of the assessment is overall moderate but can be further detailed as:

1. High in the central Baltic Sea where long-term measurements of pH and  $A_T$  exist, as well as continued  $pCO_2$  measurements recorded on the Finnmaid vessel (Figure 1 and 7).
2. Moderate in the Kattegat, the Danish Straits, and the major gulfs (i.e., the Gulf of Bothnia, Finland, and Riga, respectively), where in particular  $A_T$  has been measured more sporadically, and where  $pCO_2$  measurements are few and scattered or non-existent.

## 6 Drivers, Activities, and Pressures

The main driver of acidification is the release of anthropogenic CO<sub>2</sub>. The oceans currently absorb ~30% of our emissions (Friedlingstein et al., 2022), leading to an average pH decrease of ~0.02 per decade in the open ocean. In coastal seas, pH trends are often observed to differ largely from the open ocean trends, due to the influence of various other drivers, such as changes in eutrophication/oligotrophication and A<sub>T</sub> (Carstensen and Duarte, 2019). Deviations from open ocean pH trends are certainly observed in the Baltic Sea as well, for reasons that are not yet fully understood (Section 4.2).

A<sub>T</sub> in different areas of the Baltic Sea depends on the relative proportions of ocean water and freshwater, as well as the A<sub>T</sub> of individual rivers. Ocean water A<sub>T</sub> usually differs from river water A<sub>T</sub>, which means that linear A<sub>T</sub>-salinity relations can be observed along surface water salinity gradients (e.g. Müller et al., 2016). For that reason, changes in salinity also leads to changes in A<sub>T</sub> and in extension pH. River loads of A<sub>T</sub> can change over time both as a result of changes in precipitation and runoff, but also as a result of changing weathering rates which influence riverine A<sub>T</sub> concentrations. Surface water A<sub>T</sub> is also directly influenced by depositions of acids, such as sulphuric acid (Omstedt et al., 2015).

Changes in nutrient loads and eutrophication status can also cause significant pH trends over time (Figure 2). A more eutrophic state generally leads to larger seasonal pH variations, with higher pH maxima in summer and lower pH minima in winter (Gustafsson and Gustafsson, 2020). Pressures and activities are indicated in Table 3.

**Table 3.** Brief summary of relevant pressures and activities with relevance to the indicator.

	General	MSFD Annex III, Table 2a
<b>Strong link</b>	Pressures: <ul style="list-style-type: none"> <li>- Nutrient inputs resulting in changed eutrophication status, indirectly impacting the acidification state</li> <li>- Deposition of acidic compounds, e.g. sulfuric acid</li> <li>- Changes in hydrological conditions, e.g. A<sub>T</sub>, salinity and temperature directly impact acidification in a number of non-trivial ways</li> <li>- Change in atmospheric pCO<sub>2</sub></li> </ul> Activities: <ul style="list-style-type: none"> <li>- Activities related to nutrient inputs, e.g. agriculture, aquaculture.</li> </ul>	<u>Pressures (Table 2a)</u> <p>Physical</p> <ul style="list-style-type: none"> <li>- Changes to hydrological conditions</li> </ul> <p>Substances, litter and energy</p> <ul style="list-style-type: none"> <li>- Input of nutrients — diffuse sources, point sources, atmospheric deposition</li> <li>- Input of organic matter — diffuse sources and point sources</li> <li>- Input of other substances (e.g. synthetic substances, non-synthetic substances, radio nuclides) — diffuse sources, point sources, atmospheric deposition, acute events</li> </ul>

	<ul style="list-style-type: none"> <li>- Activities causing CO<sub>2</sub> emissions, e.g. energy generation, agriculture, transportation.</li> </ul>	<p><b><u>Activities (Table 2b)</u></b></p> <p><b>Production of energy</b></p> <ul style="list-style-type: none"> <li>- Non-renewable energy generation</li> </ul> <p><b>Cultivation of living resources</b></p> <ul style="list-style-type: none"> <li>- Aquaculture — marine, including infrastructure</li> <li>- Aquaculture — freshwater</li> <li>- Agriculture</li> <li>- Forestry</li> </ul> <p><b>Transport</b></p> <ul style="list-style-type: none"> <li>- Transport — shipping</li> <li>- Transport — air</li> <li>- Transport — land</li> </ul> <p><b>Urban and industrial uses</b></p> <ul style="list-style-type: none"> <li>- Waste treatment and disposal</li> </ul>
<b>Weak link</b>		

## 7 Climate change and other factors

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Increasing atmospheric CO<sub>2</sub> on one hand drives climate change, but is on the other hand also the main driver of large-scale and long-term pH trends – for that reason, ocean acidification is commonly referred to as “the other CO<sub>2</sub> problem” (e.g. Doney et al., 2009). Depending on emission scenario, the CO<sub>2</sub>-induced acidification in the Baltic Sea can be expected to lead to a pH decrease in a range 0.1-0.3 units by year 2100 compared to present-day values (Gustafsson and Gustafsson, 2020). But, as described in Sections 4 and 5, other processes such as eutrophication and changing A<sub>T</sub> can lead to pH trends in coastal seas that significantly differ from the open ocean acidification trend (Duarte et al., 2013; Carstensen and Duarte, 2019).

Runoff is projected to increase in the northern Baltic Sea, resulting in decreased A<sub>T</sub> and in extension pH due to the lower salinity. A higher atmospheric CO<sub>2</sub> level could on the other hand enhance weathering in Baltic Sea catchment areas, leading to increasing riverine A<sub>T</sub>. The net effect on A<sub>T</sub> is currently not known (HELCOM, 2021). Organic matter production and mineralization together with changes in temperature and stratification control seasonal surface water pCO<sub>2</sub> and pH patterns. Eutrophication leads to larger seasonal variations, i.e., higher pH maxima in summer and lower pH minima in winter, while mean pH is only affected to a smaller degree (Gustafsson and Gustafsson, 2020). Reduced nutrient loads and primary production following BSAP could thus on one hand lead to lower pH peaks in summer, but on the other hand also a less pronounced pH decrease in winter.

Nevertheless, sensitivity experiments indicate that the development of atmospheric CO<sub>2</sub> will dominate long-term pH trends in open Baltic Sea waters during the 21<sup>st</sup> century (Gustafsson and Gustafsson, 2020). The atmospheric CO<sub>2</sub> level could in a worst-case scenario exceed 950 ppm by year 2100, i.e., more than a doubling of today’s level. It is in such a case highly unlikely that the CO<sub>2</sub>-induced acidification can be counteracted by other processes in the Baltic Sea (Gustafsson and Gustafsson, 2020).

## 8 Conclusions

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Several processes influence pH and the marine carbonate system in coastal seas, but in the long run the atmospheric CO<sub>2</sub> level is likely to dominate over all other processes unless emissions are mitigated. The impact of acidification on populations of marine organisms in the Baltic Sea is so far small. In a worst-case emission scenario, detrimental effects could however be anticipated in species that cannot adapt to a warmer and more acidified Baltic Sea, and furthermore plankton community composition and food-web interactions could be altered.

It is a scientific challenge to set threshold values regarding both hypercapnia and decalcification effects (Section 3). The reason for this is the large variations between species and also within species depending on e.g. age class and adaptation. Ecosystem effects of acidification can further be strongly modified by interactions with other drivers, such as changes in temperature, salinity, stratification, and oxygen conditions. Several studies have for those reasons emphasized the need for a multi-stressor approach on community and habitat levels rather than studies of single-species response to a single stressor (e.g. Havenhand, 2011).

It is important to improve the temporal and spatial monitoring of acidification parameters to be able to identify key areas where acidification effects are likely to become a problem. In such areas, reducing the impact of other environmental stressors will become increasingly important if it is not possible to mitigate the main driver of acidification, i.e., the increasing atmospheric CO<sub>2</sub> level.

### 8.1 Future work or improvements needed

Our knowledge of the development of acidification in the Baltic Sea is currently limited to a few sites. To improve and broaden our knowledge, A<sub>T</sub> and pH should be included as standard variables in the monitoring programs of all contracting parties. In addition, the measurement procedures for pH needs to be homogenized to ensure high quality data (see also Section 9.2-3).

A<sub>T</sub> has showed increasing trends in the Baltic Sea over the past decades, but the reason is not known (Müller et al., 2016). Furthermore, A<sub>T</sub> budgets have revealed that riverine and internal A<sub>T</sub> sources are not sufficient to explain the observed A<sub>T</sub> concentrations in the Baltic Sea, indicating the existence of a significant “unresolved A<sub>T</sub> source“ (Gustafsson et al., 2019a, b). Recent findings imply that dissolution of calcium carbonate mobilized via coastal and seabed erosion can significantly contribute to A<sub>T</sub> generation, which could largely explain the unresolved A<sub>T</sub> source (Wallmann et al., 2022). This hypothesis needs further testing.

## 9 Methodology

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### 9.1 Scale of assessment

The entire Baltic Sea is covered by the assessment. Currently no assessment is made directly utilising the HELOCM Scales of assessment, but relevant sub-basins (i.e. HELCOM Scale 2 assessment units) are referred to.

### 9.2 Methodology applied

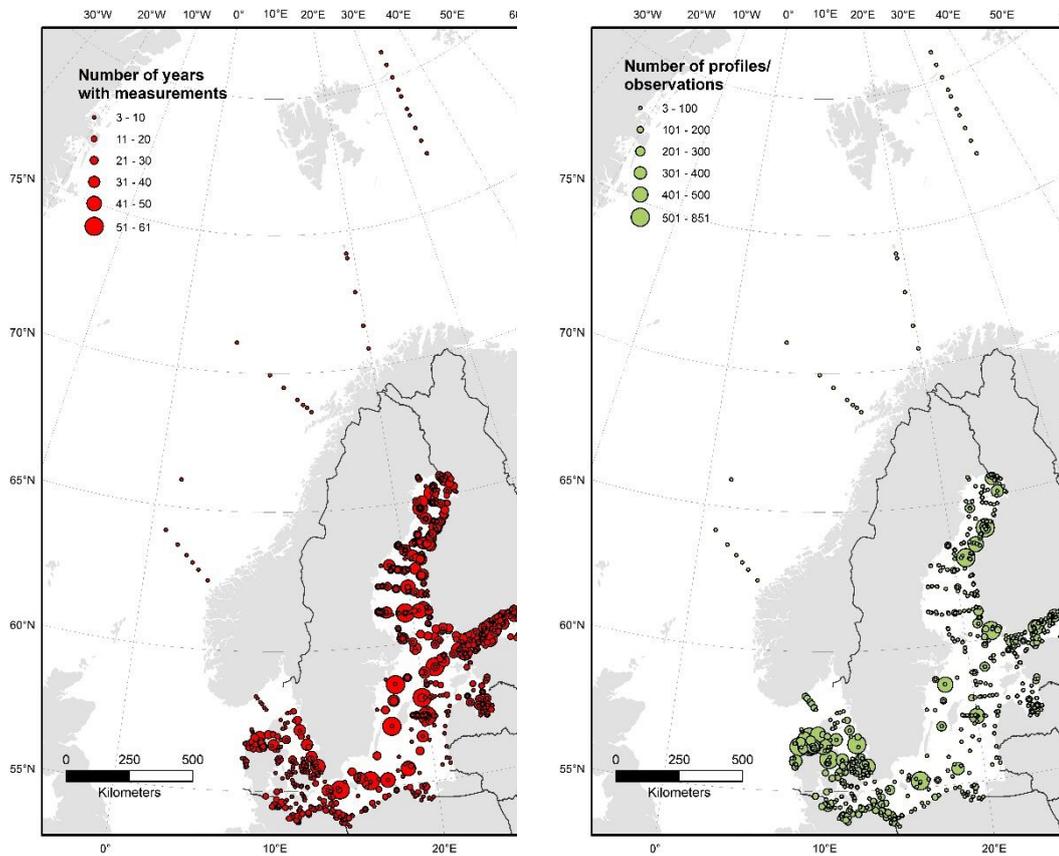
The carbonate system is described by the four parameters  $C_T$ ,  $A_T$ , pH, and  $pCO_2$ . Measurement of all four parameters is straightforward (Dickson et al., 2007).  $pCO_2$  and  $C_T$  are not explicitly mentioned in the COMBINE General Guidelines for monitoring of the Baltic Sea (HELCOM, 2006), but straightforward SOPs exist (Dickson et al., 2007; Pfeil et al., 2013), and the methodological approach does not differ from that for open ocean waters. For  $A_T$  and pH, general guidelines exist and have been updated in 2017. Yet, pH remains a challenging parameter as the generally used and still recommended potentiometric measurement would require the use of buffer solutions at ionic strength close to that of the measured sample. Also, defining appropriate standards is problematic. Recent advances in spectrophotometric pH measurements in brackish waters (Müller et al., 2018a; Douglas and Byrne 2017; Müller et al., 2018b; Müller and Rehder., 2018) suggest that this method will soon be the most accurate and precise method for pH detection also in brackish waters. At least one instrument is currently commercially available, a second to be launched soon. Thus, due to its higher accuracy and superior long-term traceability, the progress in theory and technology of this method for brackish water applications should be closely followed.

$A_T$  and pH are not consistently monitored across the Baltic Sea (Figure 8), and data have been measured and reported with varying precision and quality. Many areas only have a few spot samples, particularly in the coastal zone, whereas longer time series exist for the open Baltic Sea. The amount of pH observations increased from the 1970s and peaked in the early 2000s, reaching a contemporary level of ~1500 profiles per year (Figure 9). This is substantially lower than the similar data amount for nutrients, chlorophyll, and oxygen. A consequence is that acidification trends can be assessed only for the regular open water stations and a limited amount of coastal stations.

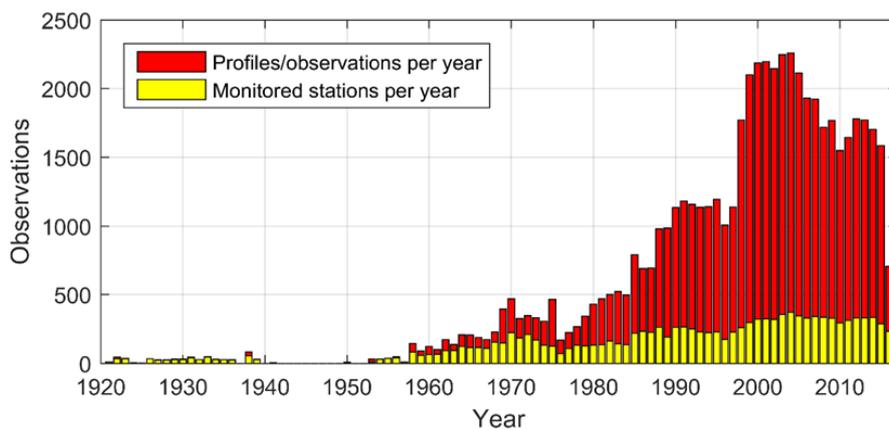
Historical data from various monitoring programs have been collected and quality assured in the Baltic Environmental Database (BED). This work has been carried out as part of a number of different research projects. Although the precision of pH measurements is generally low (typically 0.1 pH unit), trends considered are generally much larger (Figure 2) and based on many observations, which partly remedies the poor precision and accuracy of the pH data.

For the open Baltic Sea, long-term traceable monitoring of pH as the prime variable of acidification with highest possible precision and accuracy has not been achieved. One of the main reasons for this shortcoming was the lack of availability of a method for long-

term traceable measurement of pH at a resolution reflecting gradual nature in brackish water systems. Recent advances in technology now allow to overcome this problem.



**Figure 8:** Length of time series number of profiles/observations of pH across the Baltic Sea. Data from ICES and Danish monitoring program, compiled by the TRIACID project.



**Figure 9:** Amount of profiles/observations of pH over time. Data from ICES and Danish monitoring program, compiled by the TRIACID project.

### 9.3 Monitoring and reporting requirements

$A_T$  and pH should be monitored as standard hydrochemistry variables in the monitoring programs of contracting parties.  $A_T$  and pH are low-cost measurements that are fairly simple to analyse and therefore, adding these to existing monitoring programs will only marginally increase the cost of monitoring. The measurement procedures for pH should be further homogenized and mean quality of data should be improved.

## 10 Data

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The data and resulting data products (e.g. tables, figures and maps) available on the indicator web page can be used freely given that it is used appropriately and the source is cited.

## 11 Contributors

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## 12 Archive

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This version of the HELCOM core indicator report was published in April 2023:

The current version of this indicator (including as a PDF) can be found on the [HELCOM indicator web page](#).

No earlier versions of this indicator currently exist.

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## 14 Other relevant resources

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No additional components are provided in this report.